

## THE DISSIPATIVE QUASI-GEOSTROPHIC EQUATION IN SPACES ADMITTING SINGULAR SOLUTIONS\*

Yuan Baoquan

(College of Mathematics and Informatics, Henan Polytechnic University,  
Jiaozuo 454000, Henan, China)  
(E-mail: bqyuan@hpu.edu.cn)

Yuan Jia

(The Graduate School of China Academy of Engineering Physics  
P. O. Box 2101, Beijing 100088, China)  
(E-mail: sunny\_noon@sina.com)

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**Abstract** This paper studies the Cauchy problem of the dissipative quasi-geostrophic equation in pseudomeasure space  $PM^{n+1-2\alpha}(\mathbb{R}^n)$  or Lorentz space  $L^{\frac{n}{2\alpha-1},\infty}(\mathbb{R}^n)$ , which admit the singular solutions. The global well-posedness is established provided initial data  $\theta_0(x)$  are small enough in these spaces. Moreover, the asymptotic stability of solutions in pseudomeasure space is proved. In particular, if the initial data are homogeneous functions of degree  $1 - 2\alpha$ , the self-similar solutions are also obtained.

**Key Words** Dissipative quasi-geostrophic equation; singular solutions; pseudomeasure spaces; Lorentz space; global well-posedness.

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### 1. Introduction

In this paper we discuss the following Cauchy problem of the dissipative quasi-geostrophic equation:

$$\begin{cases} \frac{\partial \theta}{\partial t} + (u \cdot \nabla)\theta + \kappa(-\Delta)^\alpha \theta = 0, & (t, x) \in (0, \infty) \times \mathbb{R}^n, \\ \theta(x, 0) = \theta_0(x), \end{cases} \quad (1.1)$$

where  $0 < \alpha \leq 1$  is a fixed parameter and  $\kappa > 0$  is the dissipative coefficient. The function  $\theta(t, x)$  represents the potential temperature and the fluid velocity  $u(t, x) =$

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$(u_1, u_2, \dots, u_n)$  is divergence free and determined from  $\theta(t, x)$  by a stream function  $\psi(t, x)$ :

$$u_j = \pm \frac{\partial \psi}{\partial x_k}, \quad j, k = 1, 2, \dots, n. \tag{1.2}$$

The function  $\psi(t, x)$  satisfies

$$(-\Delta)^{\frac{1}{2}} \psi = -\theta. \tag{1.3}$$

From (1.2) and (1.3) we can obtain that

$$u_j = \pm \mathcal{R}_{\pi(j)} \theta, \quad \pi(j) \text{ is a permutation of } j = 1, 2, \dots, n, \tag{1.4}$$

where  $u_j$  may take either plus or minus sign, and

$$\mathcal{R}_j = \frac{-\partial_{x_j}}{(-\Delta)^{1/2}}, \quad j = 1, 2, \dots, n \tag{1.5}$$

are the Riesz transforms.

In the equation (1.1) the Riesz potential operator  $(-\Delta)^\alpha$  is defined through the Fourier transform

$$\widehat{(-\Delta)^\alpha f}(\xi) = |\xi|^{2\alpha} \hat{f}(\xi), \tag{1.6}$$

where  $\hat{f}(\xi) \triangleq \mathcal{F}(f)(\xi) = (2\pi)^{-\frac{n}{2}} \int_{\mathbb{R}^n} e^{-ix\xi} f(x) dx$  is the Fourier transform of  $f(x)$ . For notational convenience, we occasionally write  $\Lambda$  for  $(-\Delta)^{\frac{1}{2}}$ . In the particular case  $n = 2$ , we have

$$u(t, x) = (\partial_{x_2} (-\Delta)^{-\frac{1}{2}} \theta, -\partial_{x_1} (-\Delta)^{-\frac{1}{2}} \theta) = (-\mathcal{R}_2 \theta, \mathcal{R}_1 \theta) = \mathcal{R}^\perp \theta, \tag{1.7}$$

where  $\mathcal{R} = (\mathcal{R}_1, \mathcal{R}_2)$  is the 2-D Riesz transform.

Physically, the 2-D quasi-geostrophic equation models the evolution of temperature of atmospheric and oceanic fluid flow on the two dimensional boundary of a fast rotating three dimensional half space with small Rossby and Ekman number [1, 2]. The scalar  $\theta(t, x)$  represents the potential temperature and  $u(t, x)$  is the fluid velocity. These equations have been actively investigated because of the mathematical importance and potential applications to meteorology and oceanography [1–4].

Mathematically, the 2-D quasi-geostrophic equation serves as a lower dimensional models of the 3-D Navier-Stokes equations because of the striking similarity between the behavior of its solutions and that of the potential singular solutions of the 3-D N-S equations [1], thus the study of the 2-D quasi-geostrophic equation may provide some clues to the millennium prize problems on the Navier-Stokes equations [5].

In a series of previous works of Wu [6–9], the well-posedness results for initial data  $\theta_0$  in Lebesgue space  $L^p$ , homogeneous Sobolev space  $\dot{L}_{s,p}$ , Morrey space  $M_{p,\lambda}$  and Hölder